

Composition and metabolism of microbial communities in soil pores

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Delineation of microbial habitats within the soil matrix and characterization of their environments and metabolic processes are crucial to understand soil functioning, yet their experimental identification remains persistently limited. We combined single- and triple-energy X-ray computed microtomography with pore specific allocation of ¹³C labeled glucose and subsequent stable isotope probing to demonstrate how long-term disparities in vegetation history modify spatial distribution patterns of soil pore and particulate organic matter drivers of microbial habitats, and to probe bacterial communities populating such habitats. Here we show striking differences between large (30–150 μm Ø) and small (4–10 μm Ø) soil pores in (i) microbial diversity, composition, and life-strategies, (ii) responses to added substrate, (iii) metabolic pathways, and (iv) the processing and fate of labile C. We propose a microbial habitat classification concept based on biogeochemical mechanisms and localization of soil processes and also suggests interventions to mitigate the environmental consequences of agricultural management.

Soil is a crucial component of all terrestrial ecosystems, the main source of human food production, and one of the major mediators of atmospheric CO₂ level^{1,2}. Soil functioning is enabled by a complex community of microorganisms³ inhabiting an intricate physical frame of soil pores^{4,5}, adapting their life and C-acquisition strategies⁶ to highly variable micro-environmental conditions^{7,8}.

Plant roots are among the main drivers of soil micro-environments^{9,10}, and vegetation community composition and diversity substantially affect formation of microhabitats within the soil matrix¹¹. Plant roots form and modify soil pore structure, i.e., pore-size distributions, pore shapes and surface properties, and pore connectivity¹², while pore structure in-turn impacts soil microorganisms^{13–15}. Greater plant species richness increases root biomass and diversifies root residue inputs¹⁶, providing more organic C for soil microorganisms¹⁷ and raising microbial activity^{17–19}. Plant communities, especially the diverse perennials, strongly influence soil

pores—form biopores, leading to structural changes positively associated with soil organic matter accumulation²⁰.

The pore structure and distribution as well as fluxes of water within pores ultimately define oxygen and nutrient supply for microbial functioning^{21,22}. Pores in the <10 μm Ø size range are often saturated by water, which limits oxygen supply, while pores in the >1000 μm Ø are mostly unsaturated and even dry²³. Pores with diameters in the few-tens to couple-hundred-micron size range provide an optimal balance of oxygen, water, carbon (C), and nutrient inflows for resident microorganisms. Therefore, microbial communities inhabiting pores of various sizes differ in their composition, life strategies, and activities: >30 μm Ø pores can better stimulate fast decomposition of newly added C and have a greater abundance of certain microbial taxa, as compared to <10 μm Ø pores^{24–26}. Yet, even this is a simplification, subject to actual proximity to substrates, connectivity, and the presence and accessibility of nearby water²⁷. Water as

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films and menisci between soil particles provide habitable niches and crevasses for microorganisms to grow, function, move, and interact^{28–31}.

Thus the movement of soil solution, and, in particular, pore-scale hydraulic connectivity, is a key element influencing bacteria and especially their access to organic matter and subsequent metabolism^{32–34}, with ultimate consequences for soil carbon accretion and stability and for ecosystem processes such as decomposition and denitrification^{22,35}. Yet we know very little about the impact of hydraulic connectivity on microbes at the scale of individual soil pores. While the expansion of X-ray computed microtomography (μ CT) has made it possible to assess and characterize pores in intact soil cores^{36,37}, the visualization of water has been much more challenging and therefore, not yet combined with pore-size visualization for assessing the full impact of pore structure on ecosystem processes.

Here we document the combined effects of long-term differences in vegetation history on the spatial distribution of soil pores and particulate organic matter (POM) as well as on the hydraulic connectivity of small (4–10 μm \emptyset) and large (30–150 μm \emptyset) pores within intact soil matrices. We further test the degree to which vegetation history affects bacterial richness, community composition, and metabolism, specifically in small (4–10 μm \emptyset) vs. large (30–150 μm \emptyset) pores within the soil matrix, with consequent effects on soil C processing. We used three vegetation systems of contrasting managements and plant diversities established in a replicated blocked field experiment: (i) a multiyear fallow followed by 2 years of monoculture corn (*Zea mays* L.), which for brevity we will refer to as bare soil, (ii) a perennial monoculture switchgrass (*Panicum virgatum* L.) community, and (iii) a polyculture restored prairie community of >20 native North American grasses and forbs. We used μ CT to characterize soil pore structure and localize and quantify the spatial distribution of particular organic matter (POM), employing a triple-energy μ CT approach to examine pore-level spatial patterns of water distributions. We simulated labile substrate additions by applying small quantities of labeled glucose—an abundant component of root exudates and decomposition product of carbohydrates that is metabolized by the majority of soil microorganisms with well-defined uptake mechanisms and mineralization pathways³⁸. We then used stable isotope probing (¹³C-DNA/RNA-SIP)^{39–41} to identify microorganisms actively assimilating the glucose-derived C (¹³C) into nucleic acids^{42–44}. Results demonstrate that microbial habitats defined by soil pores and POM of various plant communities differ in ways that strongly influence microbial composition and activity, and thus may impact ecosystem processes such as decomposition, nitrogen processing, and carbon sequestration.

Results

Soil chemical and physical characteristics

The soil under restored prairie developed higher C and N contents, lower bulk density, and larger total porosity than the switchgrass soil (Supplementary Table S2). The average distance to pores tended to be the lowest in the prairie soil (Fig. 1c and Supplementary Table S2), suggesting that the distribution of pores through the matrix of the prairie soil was the most uniform of the three systems.

Large (30–150 μm \emptyset) and small (4–10 μm \emptyset) pores in soils of all vegetation systems markedly differed in their connectivity. The large pores not connected to the main pore space constituted 0.1% of the total soil volume (Fig. 1d), i.e., amounted to just a small fraction of all large pores (3.8–5.3%, Supplementary Table S2). In contrast, out of the total 1.5–1.9% of the soil volume occupied by small pores (Supplementary Table S2), >60% were disconnected from the main connected pore space (Fig. 1d).

The two perennial vegetation systems, i.e., switchgrass and prairie, contained approximately five times more POM than the bare soil system (Supplementary Table S2). The average distance to POM fragments in the prairie soil was only half of that in the switchgrass

(Fig. 1c and Supplementary Table S2), indicating a much more homogeneous distribution of POM within the prairie soil's matrix. The size of POM fragments in the prairie soil was, on average, smaller by one-third as compared to switchgrass (Fig. 1e), and the prairie soil had >5 times more POM fragments than the switchgrass soil (Fig. 1f). The high content of uniformly distributed POM fragments resulted in most of the total volume of the prairie soil (i.e., 87% of it) to be in close proximity (<300 μm) to and, thus, under the direct influence of POM. Only 48% and 30% of the soil volumes of the switchgrass and bare soils, respectively, were within <300 μm distance to POM (Fig. 1g).

Dopants and glucose additions to small and large pores

Analyses of the multi-energy images from μ CT scans revealed that the dopant solutions reached the target pores (Fig. 1b). The small (4–10 μm \emptyset) pores were filled by the added KI solution to ~58% saturation; the saturation of the large (30–150 μm \emptyset) pores by BaCl_2 was 10–15% (Supplementary Fig. S2). Given the high variability and connectedness of the soil pore space, we were not surprised that some portions of the solutions were found in the pores of intermediate size (10–30 μm \emptyset), and there was some small overlap in saturation between the small and large pores. However, the liquid intended for the large pores did primarily occupy pores >50 μm \emptyset , with a peak saturation around 100–150 μm , which was within the target size.

The solution added to the large pores formed sizeable menisci between soil particles as well as thick water films on the boundaries of very large pores, which had their central portions filled with air (Fig. 1b). The space filled with the liquid intended for the large pores was much better connected than that filled with the solution added to the small pores. Specifically, the largest connected pore volume filled with large-pore-targeting solution occupied 2.5% of the total μ CT-visible pore space (>4 μm \emptyset pores). The largest connected pore volume filled with small-pore-targeting solution occupied only 0.4% of the pore space.

Glucose originated carbon

The switchgrass and prairie systems contrasted in the temporal dynamics of ¹³C-CO₂: on the 1st day of the incubation, the ¹³C enrichment was higher in the CO₂ released from the prairie soil, while during the subsequent few days (days 3 and 7) the ¹³C enrichment was higher in the switchgrass (Fig. 2a and Supplementary Table S3). On the first day after glucose input, ¹³C atom-% of the CO₂ released from the large pores of switchgrass soil tended to be lower than that from small pores.

After the first 24 h of glucose utilization, the total ¹³C amounts remaining in the large and small pores were similar (Fig. 2b). Approximately 45% of the total ¹³C recovered within the switchgrass soil after 24 h was present as DOC (Fig. 2c), while within the prairie soil it was only ~18%. Using a conversion factor of 0.45 for the fumigation approach to microbial biomass determination, the ¹³C incorporation into microbial biomass constituted on average ~33% of the ¹³C remaining in the soil (Fig. 2d).

After the 30-day incubation, the ¹³C remaining in the soil was higher in the small compared to the large pores (Supplementary Table S3), with the difference especially pronounced in the switchgrass soil. The dissolved organic ¹³C constituted only ~2% of the total ¹³C recovered in the prairie soil, and was even lower (<1%) in the small pores of the switchgrass cores. ¹³C of microbial biomass constituted ~20% and 12% of the total soil ¹³C in the prairie and switchgrass soils, respectively.

Microbial community analysis

The perennial vegetation was a major determinant of the microbial community composition (Supplementary Figs. S3 and S4). Compared to the bare soil, the relative abundance of phyla Latescibacteria, Gemmatimonadetes, Planctomycetes, Proteobacteria, and

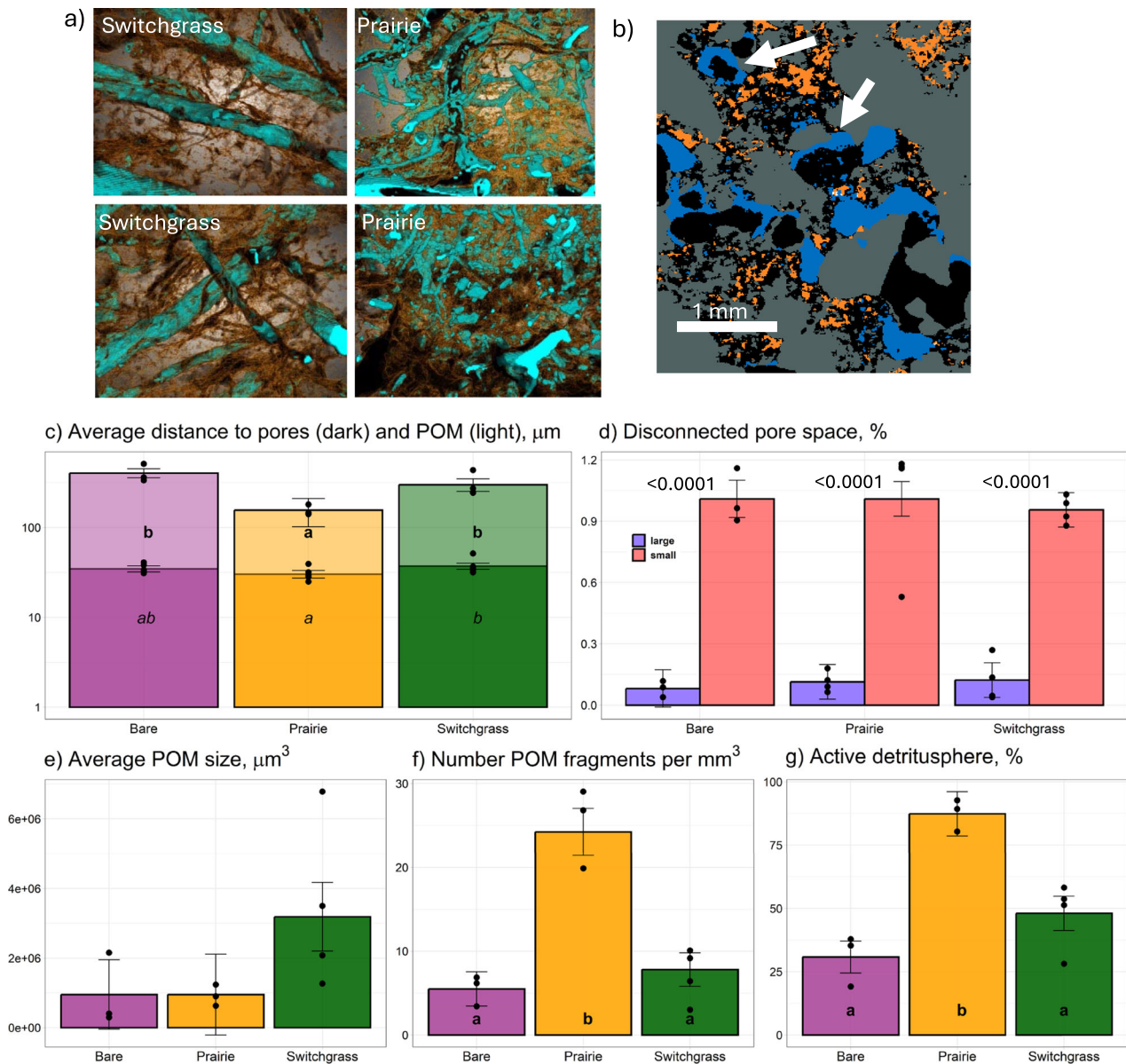


Fig. 1 | Images and characteristics of soil pores $>4 \mu\text{m} \text{Ø}$ and particulate organic matter (POM) within the soils of three experimental plant communities.

a Sample images with pores (brown) and POM (cyan) from selected representative soil cores of prairie and switchgrass soils. **b** Visualization of the liquid added to small (4–10 $\mu\text{m} \text{Ø}$) (orange) and large (30–150 $\mu\text{m} \text{Ø}$) (blue) pores. White arrows are pointing to examples of menisci between soil particles and water films within the large pores. **c** Average distances to pores (dark shading) and POM (light shading) in the soils of the three communities. **d** Volumes of large and small pores that were not connected to the largest connected pore cluster (% of total soil volume).

e Average size of POM fragments. **f** Number of POM fragments per unit of soil volume. **g** The percent of the total solid volume that was within 300 μm of POM, and thus considered to be active detritusphere. For all panels, shown are means, standard errors as error bars, and original data as dots. The letters on (c, f, g) mark the differences among the vegetation systems significant at $P < 0.05$ (bold) or $P < 0.1$ (italic). The results of the two-sided tests for the differences between the large and small pores are reported with P values shown above bars on (d). Source data for (c–g) are provided as a Source Data file.

Verrucomicrobia (Supplementary Fig. S4) was greater in the prairie and switchgrass systems. In contrast, the relative abundances of oligotrophic (*Acidobacteriales*), slow-growing (*Frankiales*), pseudomycelium-forming (*Actinomycetota*), spore-forming (*Ktedonobacteriales*), and desiccation-resistant (*Firmicutes*) bacteria were higher in bare versus planted soil. Lower pH of the bare soil might have been a contributor to a greater abundance of *Acidothermus*.

All diversity indices pointed to a lower microbial community diversity in the bare soil as compared to the prairie and switchgrass soils (Fig. 3a and Supplementary Table S4). Across both incubation times, the microbial diversity was higher in the prairie than in switchgrass soil, and at the end of the incubation, it was higher in the small

than in the large pores (Fig. 3a, $P < 0.05$). Microbial biomass was higher in the prairie soil than in the soils of bare and switchgrass systems (Supplementary Table S2).

¹³C-enriched phylotypes: immediate response to glucose addition (24-h incubation)

The total number of OTUs that responded to ¹³C-glucose within 24 h was substantially greater in the switchgrass and bare soil as compared with the prairie soil (64 OTUs) (Fig. 3b). *Pseudomonas* was the group with the highest ¹³C enrichment in the large pores of all three systems, with the ¹³C enrichment in the prairie soil twice as high as that in any other group (Fig. 3c and Supplementary Table S5). There were several

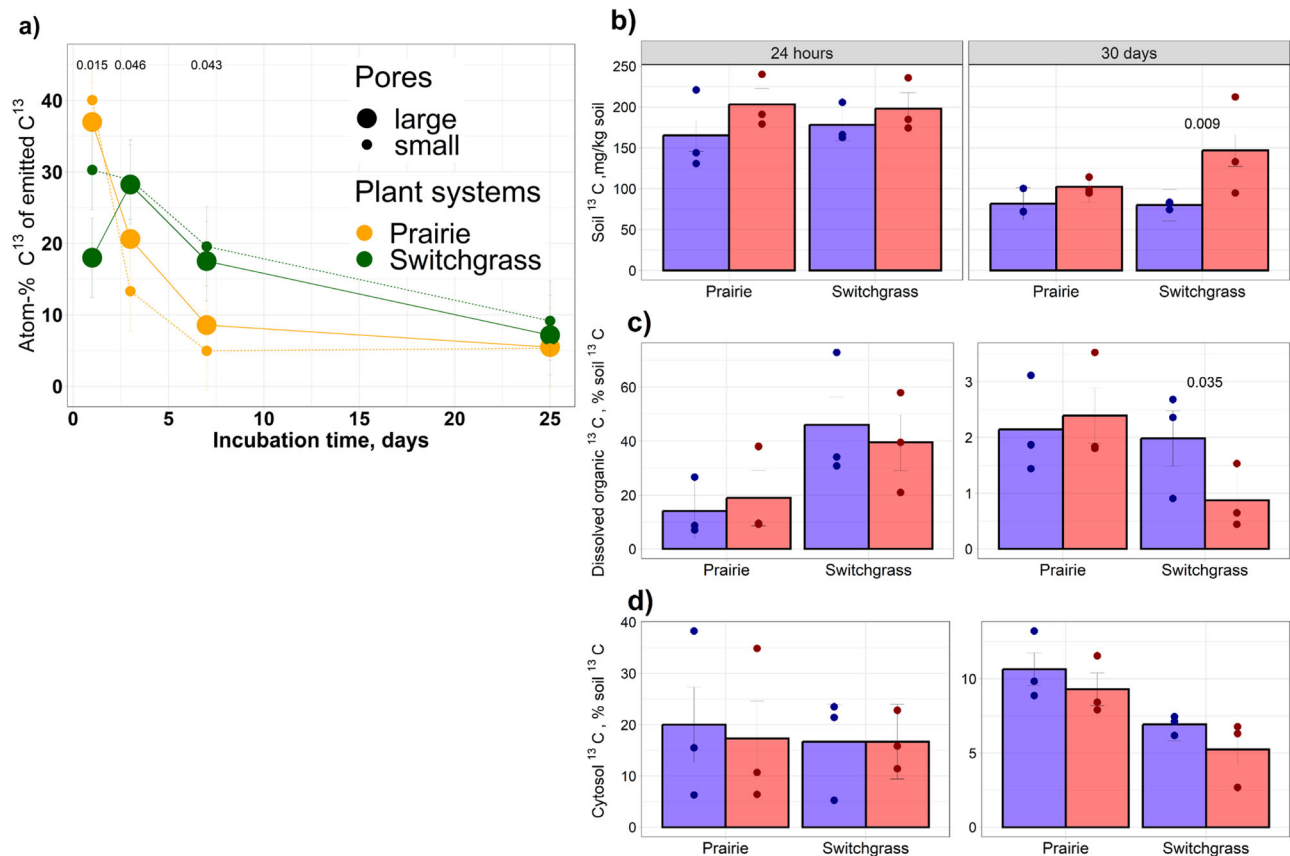


Fig. 2 | Fate of ^{13}C -glucose carbon added to small (4–10 μm \varnothing) and large (30–150 μm \varnothing) pores of the prairie and monoculture switchgrass communities. a Atom-% ^{13}C of emitted CO_2 during the 30-day incubation of the intact samples. Shown are means and standard errors as error bars. P values are shown for the differences between the two systems significant at $P < 0.05$. Total (b), dissolved organic (c), and cytosol (d) C of glucose-origin (^{13}C) remaining in the soil after 24 h

and 30-day incubation. Shown are means, standard errors as error bars, and original data as dots. Significant ($P < 0.05$) differences between large and small-pore treatments within the system and incubation times as determined via simple-effect F tests are marked, with P values shown above the bars. Source data are provided as a Source Data file.

groups of highly ^{13}C -enriched bacteria that were plant system-specific (Fig. 3c and Supplementary Table S5). The taxa *Planococcaceae* and *Micrococcaceae* incorporated ^{13}C only in the bare soil, while *Cupriavidus* and *Duganella* incorporated ^{13}C only in the switchgrass soil. *Clostridium* took the ^{13}C in the large and, especially, small pores of prairie and switchgrass.

^{13}C -enriched phylotypes: short-term response to glucose addition (30-day incubation)

In contrast to the immediate response to the glucose input, the total number of OTUs enriched with ^{13}C after a 30-day incubation was almost three times larger in the prairie (259 OTUs) than in the switchgrass soil (90 OTUs) (Fig. 3b). While only few taxa remained ^{13}C -enriched or even increased their ^{13}C enrichment with time, several new groups became enriched in all pores and vegetation systems (Fig. 3c and Supplementary Table S6). *Cupriavidus* and *Duganella* remained enriched in the large and small pores of the switchgrass soil, while absent in the prairie. *Bradyrhizobium* and *Bdellovibrio* were among the top-enriched bacteria found only in the prairie but not in the switchgrass soil (Supplementary Table S6).

Approximately 75% of the ^{13}C -enriched bacteria were present in both large and small pores in the prairie soil, while only 48% were present in both large and small pores in the switchgrass soil (Fig. 3b). *Cellvibrio* and *Bacteriovorax* were among the organisms remaining ^{13}C -enriched in the large pores (Supplementary Table S7). *Pseudomonas* remained the most (prairie) or the second most (switchgrass) enriched bacteria in the large pores, yet its ^{13}C enrichment was much

lower in the small pores of both systems (Supplementary Table S6). Several members of *Bdellovibrio* group were enriched in the large pores of the prairie soil, while this group had only very low ^{13}C incorporation in the small pores (Supplementary Table S6).

^{13}C -enriched functional genes: glycolysis/gluconeogenesis and citric acid cycle

The lack of ^{13}C enrichment in genes responsible for the first step of glycolysis in the large pores of the switchgrass system ([EC:2.7.1.2], [EC:2.7.1.63]) was noted in both 24 h and 30-day incubations, while the relevant genes were enriched in the small pores (Supplementary Fig. S5). On the contrary, the ^{13}C enrichment was common in large and small pores of the prairie soil. In the 24 h incubation, the genes encoding pyruvate dehydrogenase ([EC:1.2.4.1], [EC:2.3.1.12]), the enzyme responsible for conversion of pyruvate to acetyl-CoA, were ^{13}C -enriched in the switchgrass but not in the other two systems.

In the prairie soil, almost all the genes involved in glucose-to-pyruvate conversion steps of glycolysis were enriched in both small and large pores at both incubation times. In the switchgrass system, the genes responsible for glucokinase production were enriched only in the small-pore treatment in both the 24 h and 30-day incubation. The genes encoding the enzymes involved in dihydrolipoamide dehydrogenase [EC:1.8.1.4] were activated in the small pores of bare soil in the 24-h incubation (Supplementary Figs. S5a and S6a). In the 30-day incubation, the genes encoding the enzymes involved in pyruvate metabolism (pyruvate dehydrogenase, dihydrolipoamide dehydrogenase [EC:1.8.1.4]) were enriched only in the small pores of both

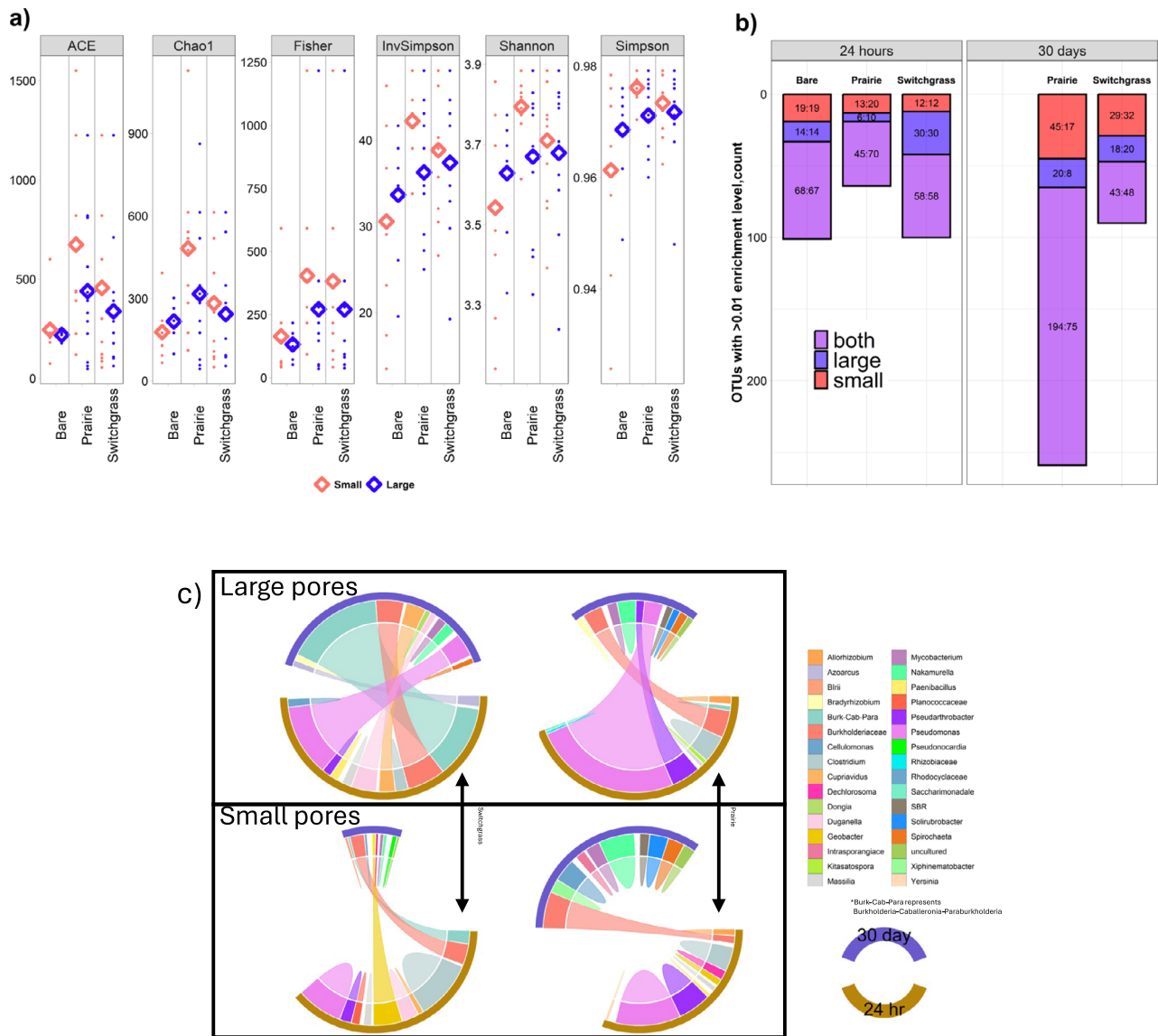


Fig. 3 | Microbial community characteristics after 24 h and 30 d of incubation in the soils of the three plant communities, with glucose placed either in large or in small pores. a Alpha diversity assessed by richness (Chao1, ACE) and diversity (Shannon, Simpson, Inverse of Simpson and Fisher) indexes. Note: the hollow diamonds represent the means of the indices in control, large pore, and small-pore incubation treatments, averaged across both incubation times for prairie and switchgrass. **b** Numbers of operational taxonomic units (OTUs) that responded to

glucose addition that were found either only in large, small, or in both pore-size treatments. Shown on each bar are OTU counts (first number) and percent of the total number of OTUs (second number). **c** Top ten genera that were ¹³C-enriched in the prairie and switchgrass plant systems in 24 h and 30-day incubations. The width of the base for each genus is proportional to its relative enrichment level. Source data for (a, b) are provided as a Source Data file.

switchgrass and prairie systems (Supplementary Fig. S6b). While in the 24-h incubation, the genes responsible for pyruvate fermentation to ethanol via alcohol dehydrogenase ([EC:1.1.1.1], [EC:1.1.2.8]), an anaerobic metabolism, were enriched both in the large and small pores of the switchgrass, after the 30-day incubation they were only enriched in the switchgrass’s small pores.

¹³C-enriched functional genes: nitrogen metabolism

As indicated by the ¹³C enrichment of the nirK gene, the bacteria benefiting from added glucose during the first 24 h of incubation performed denitrification in the switchgrass, but not in any other systems (Supplementary Fig. S7). While later (after 30 days), several genes involved in coding denitrification enzymes, namely, napA, nirK, and nosZ, were enriched in the prairie, but not in the switchgrass system.

Genes encoding carbonic anhydrase ([EC:4.2.1.1]), the enzyme that converts CO₂ and water into carbonic acid, and nitrite reductases ([EC:1.7.1.15], [EC:1.7.2.2]), the enzymes involved in dissimilatory nitrate reduction, were enriched in small pores of both prairie and switchgrass. Small pores of the switchgrass system had enrichment of the gene encoding nitrate/nitrite transport complex (nrtA) in the 24 h of incubation (Supplementary Fig. S7a) and also of the gene involved in N fixation nitrogenase (nifG) at both incubation times (Supplementary Figs. S7a and S7b). None of these genes were active in any other vegetation systems.

In the 24-h incubation, in both prairie and switchgrass and in the 30-day incubation in the prairie, the CO₂-fixing gene [EC:4.2.1.1] was ¹³C-enriched only in the large pores (Supplementary Figs. S7a and S7b).

In the 30-day incubation, glutamate dehydrogenase genes ([EC:1.4.1.2], [EC:1.4.1.4]) were enriched only in the large pores of both

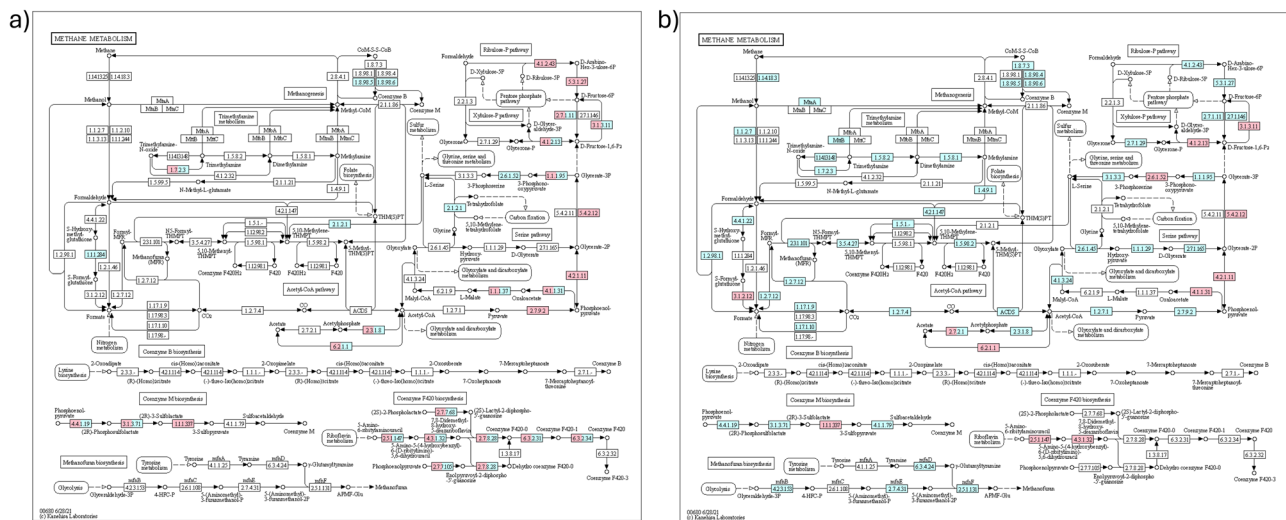


Fig. 4 | Methane metabolism genes enriched in the soil of the studied systems after 30-day incubations. Shown are the genes enriched in the prairie (a) and switchgrass (b) communities. Blue and pink denote enrichments in small and large pores, respectively. Diagrams created with KEGG (Copyright Permission 230979).

systems. Yet, glutamate synthesis ([EC:1.4.1.13], [EC:1.4.7.1]) was taking place in large and small pores of both systems.

¹³C-enriched functional genes: methane metabolism

After 24 h and 30 days, there were many enriched genes associated with the methane metabolism in the switchgrass system as compared to the prairie (Figs. 4 and S8). In switchgrass soil, gene encoding anaerobic carbon-monoxide dehydrogenase (E.1.2.7.4) was enriched in the large pores at 24-h incubation and in the small pores after 30 days. The glucose localization effects differed between the three soil–plant systems and overall, methane metabolism mainly occurred in the switchgrass and in the small pores (Fig. 4).

Discussion

Multiple years of dissimilar vegetation not only affected soil C and N contents but led to contrasting micro-scale patterns in the spatial distribution of soil pores and organic residues. As expected, many bacterial generalist taxa commonly found in soils worldwide^{45,46} were also detected in the soils of the studied systems. Yet the microbial groups residing in large vs. small soil pores responded differently to the resource addition. The responses differed both in terms of life strategies and in the metabolic pathways they employed, with implications for the fate and protection of the added glucose, and, as inferred from genetic analysis, also for other metabolic processes such as denitrification and methanogenesis.

Differences in POM spatial distributions (Fig. 1c, e, f) might be the key driver of the reported differences in decomposition between soils of monoculture switchgrass and prairie⁴⁷. A ubiquitous spread of small POM fragments in the prairie soil positioned most of the prairie soil's matrix within an active detritosphere, i.e., <300 μm distance from POM. Since many of such POM fragments were former roots, the 1–2 mm areas around them were also previously within an active rhizosphere. The detritosphere experienced major C and nutrient influxes from the decomposing organic residues^{48,49} and, in the prairie soils, acted as one giant contiguous hotspot of microbial activity⁵⁰. The large sizes of POM pieces in switchgrass (~3.5 × 10⁶ μm³) as opposed to prairie (~0.9 × 10⁵ μm³) (Fig. 1a, e), and large average distances to POM in switchgrass (Fig. 1c), suggest that a substantial portion of the switchgrass soil matrix was not under the direct influence of either detritosphere or former rhizosphere, thus, was likely devoid of fresh nutrient inputs. Likewise, most of the matrix in the bare soil was neither in the current detritosphere nor former rhizosphere. This result is especially notable because POM comparisons among plant

communities commonly focus only on total POM contents rather than on spatial distribution patterns⁵¹. Our findings (Supplementary Table S2) demonstrate that even though soils from switchgrass and restored prairie systems can have similar POM contents (e.g., ref. 52), their POM spatial distribution patterns can differ substantially, generating diverging impacts on soil C processing and microbial communities.

Microbial response to new substrate and implications for C processing differed in the immediate (i.e., 24 h) and longer (30 d) duration. Initially, i.e., 24 h after the application, the glucose addition approach equalized conditions within the micro-environments to which the glucose was added. The same quantities and concentrations of the glucose solutions, which were made using dH₂O with high oxygen levels, were added to both large and small pores. Moreover, the oxygen/glucose-rich solution was pushed into the small pores, which does not happen under natural conditions. We surmise that the resultant burst of microbial activity in both large and small pores probably led to similarly sizeable collapses in O₂ concentrations. This explanation is supported by (1) the abundance of anaerobic organisms among the initially enriched taxa, e.g., *Clostridium* (Supplementary Table S5), enrichment in several genes responsible for anaerobic processes, including the anaerobic pyruvate to ethanol pathway in both large and small pores (Supplementary Fig. S5a), and anaerobic carbon-monoxide dehydrogenase activity in the large pores of the switchgrass (Supplementary Fig. S8). Not surprisingly, the same newly added C source and the same resultant environmental conditions within 24 h of glucose addition led to the same groups, i.e., *Pseudomonas*, *Burkholderiaceae*, *Pseudarthrobacter*, and *Clostridium*, being among the top glucose consumers, heavily enriched in both large and small pores of bare, prairie, and switchgrass soils (Supplementary Table S5).

Despite these similarities, inherent differences in microbial community size and composition of soils under the three vegetation systems led to differences in processing new C. Over the years of bare soil management and monoculture switchgrass cultivation, the microbial communities apparently became less diverse as they adapted either to overall limitations in new inputs in bare soil or to a narrow range of C sources of switchgrass roots and rhizodeposition in the switchgrass system⁵³. In contrast, the high plant diversity of the prairie system led to a greater diversity in chemical composition and types of root inputs⁵⁴ and to a greater heterogeneity of the pore space, together maximizing microbial diversity in its small pores (Fig. 3a). Thus, as can be inferred from low ¹³CO₂ emissions and high dissolved organic ¹³C

concentrations in the switchgrass soil at the start of the incubation (Fig. 2a), the microbes there were delayed in utilizing the new resource. In contrast, the diverse and large microbial community of the prairie soil immediately reacted to glucose addition. The glucokinase activity in the prairie soil might have given a competitive advantage for some bacterial groups in glucose consumption (Supplementary Figs. S5 and S6). Lack of such activity from the inhabitants of the large pores of the switchgrass soil is probably what resulted in a slow utilization of the newly added glucose.

After 30 days of incubation the micro-environmental conditions in the soil matrix likely recovered to their original natural state. That is, after the initial oxygen depletion in the large pores of both switchgrass and prairie systems, the exchange with the atmosphere would have restored oxygen concentrations, allowing aerobic microbial groups to flourish. The oxygen-tolerant initial top consumers of ^{13}C -glucose remained highly ^{13}C -enriched until day 30, which reflected their slow turnover and low cross-feeding intensity in the community (Fig. 3c and Supplementary Tables S5 and S6).

In the small pores of the switchgrass soil, the micro-environmental conditions changed substantially from immediately after the glucose addition, bringing marked shifts in the community composition of ^{13}C -enriched organisms. We surmise that because of the poor contact with the atmosphere the initial oxygen depletion due to active glucose oxidation was not relieved in the small pores. While the glucose resource disappeared the anoxic conditions still proliferated, prompting processes such as anaerobic fermentation of pyruvate to ethanol (Supplementary Fig. S5b), dissimilatory nitrate reduction (Supplementary Fig. S7b), and methanogenesis (Fig. 4). The remaining necromass apparently was poorly utilized by the organisms that succeeded the first glucose responders (Fig. 3c). The number of these secondary ^{13}C consumers decreased and their ^{13}C enrichment dropped as compared to those at 24 h after glucose input (Fig. 3c). It seems probable that the microbial community of the small pores in switchgrass soil shifted to dormancy or basic maintenance after glucose was exhausted, as opposed to active successional development and cross-feeding.

In the prairie soil, more than 100 previously non-enriched bacterial groups now acquired the ^{13}C label in both large and small pores (Fig. 3b). The extracellular metabolites and necromass of the first responding community were heavily utilized by a very large and diverse group of successive organisms (Fig. 3b, c and Supplementary Table S6). More of the added label was still embedded in living microbial biomass or was in a form of DOC in the soil solution, while less of it was part of the ^{13}C remaining in soil as compared with small pores of switchgrass (Fig. 2b–d).

A possible explanation for the similarity between the large and small pores in the prairie soil is that most of the soil matrix harboring small pores belonged to the detritosphere (Fig. 1f and Supplementary Table S2). Thus, regardless of whether glucose was added to large or small pores in the prairie soil, it was consumed by the detritosphere microorganisms, well adapted to benefit from the labile resources both immediately (Fig. 2d) and in the medium-term. In contrast to the switchgrass system, the community of secondary decomposers in small pores of prairie remained as highly enriched after the 30-day incubation as they were 24 h after glucose input (Fig. 3c). Intensive microbial turnover resulted in strong successional changes in the community, yet the extra resource (i.e., C from the added glucose) was re-utilized by ever larger groups of organisms.

Several processes involved in N metabolism tended to be system- and pore-specific (Supplementary Fig. S7), suggesting differences in N availability and the adaptations of resident microbial communities. We surmise that N availability was low in the small pores, especially in the switchgrass soil, prompting a greater diversity of routes of N procurement by their bacterial inhabitants. For example, small pores of the switchgrass soil were populated by (i) bacteria with fast N uptake,

as suggested by an enrichment of the gene coding for nitrate/nitrite transport complex (nrtA) in the 24-h incubation, and (ii) bacteria involved in N fixation, as suggested by the gene involved in N fixation nitrogenase (nifG) in both incubation times. On the other hand, the glutamate dehydrogenase pathway, which is utilized for glutamate synthesis under conditions of N excess⁵⁵, was active in the large, but not small, pores of both systems. The genes involved in dissimilatory nitrate reduction (DNR) were enriched in the small pores of both prairie and switchgrass systems (Supplementary Fig. S7b), while denitrification-related genes were enriched only in the prairie system. This result is also suggestive of greater N deficiency in the small pores of the switchgrass system as compared with prairie, because DNR occurs at higher C:N ratios than denitrification^{56,57}.

Small-pore environments of the switchgrass and bare soil were also deficient of readily available C, likely stimulating the microorganisms to employ a variety of frugal C-use strategies. The dominance of genes associated with the later steps in the glycolysis/gluconeogenesis pathways, e.g., dihydrolopoamide dehydrogenase as activated in the small pores of the switchgrass in the 30-day incubation (Supplementary Figs. S5b and S6b) and in bare soil in the 24-h incubation (Supplementary Figs. S5a and S6a), suggests that residents of the small pores were more likely to consume the ^{13}C labeled products of initial utilization, supplementing Acetyl-CoA (Supplementary Fig. S5b) and Succinyl-CoA (Supplementary Fig. S6b). Enriched genes for carbon-monoxide dehydrogenase suggest that the organisms were striving to obtain extra C resources (e.g., CO_2 and CO), helping them to survive in an inhospitable environment^{58,59}.

Conceptual model of soil microhabitats

The observed responses of bacteria taxa to glucose additions in pores of contrasting sizes enable us to (i) postulate the delineation of three distinct micro-habitat types within the soil matrix (Fig. 5) and (ii) conceptualize their key characteristics (Table 1). The micro-habitat types are:

1. large-pore (Lp) habitats, which are spatially well-connected, with prevailing oxic conditions and root-originated C sources;
2. substrate-rich small-pore (SpRich) habitats, with somewhat restricted hydraulic connectivity and oxygen availability yet with abundant supplies of C and nutrients; and
3. substrate-poor small-pore (SpPoor) habitats, with restricted connectivity, deficient in oxygen, nutrients, and C.

In our experiment, the Lp habitat is represented by the large pores containing plant roots that release rhizodeposits including easily available exudates^{60,61}. High pore connectivity (Fig. 1d) provides more O_2 as compared to the small pores³¹, and the high hydraulic connectivity (Fig. 1b) suggests greater opportunities to receive labile substrates and nutrients carried in by frequent and fast water fluxes^{62,63}. Therefore, microorganisms in the large pores are accustomed to periodic inputs of labile C, including glucose in our experiment. The large pores would be even more subject to episodic glucose additions. Indeed, to protect their membranes under water shortage stress, microorganisms synthesize trehalose⁶⁴, which upon rewetting hydrolyzes into glucose^{65,66}. This process is particularly relevant for large pores, since they experience greater variations in water regimes, i.e., droughts followed by quick rewetting, than do small pores⁶⁷.

The SpRich habitat is represented by the small pores of the prairie soil, vast majority of which were in the active detritosphere (Supplementary Table S2), while the SpPoor, by the small pores of switchgrass and bare soils. We surmise that the proximity to roots and POM separates the small pores into SpRich and SpPoor habitats. The SpRich habitats are, in essence, located in rhizosphere and detritosphere. When in a rhizosphere, SpRich habitats receive DOC inputs from nearby roots, mycorrhiza, and root hairs, while in the detritosphere SpRich habitats receive DOC from decomposing POM. Due to their location in relative proximity to roots and larger pores, SpRich habitats

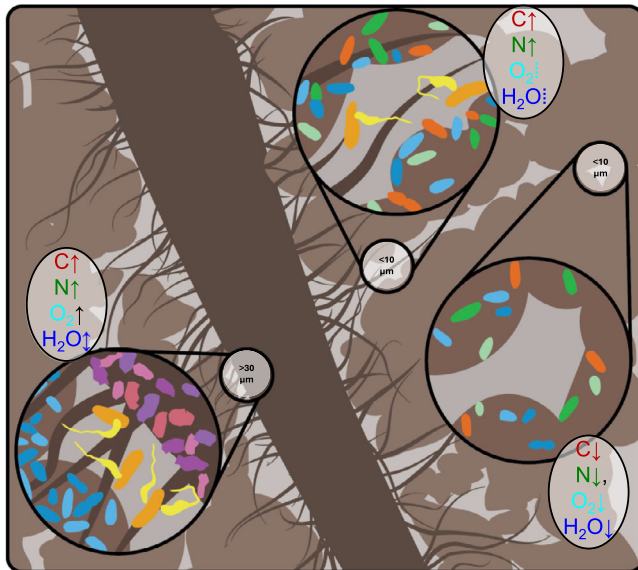


Fig. 5 | Classification concept of soil microhabitats: large pores (Lp), substrate-rich small pores (SpRich), and substrate-poor small pores (SpPoor). (1) The Lp habitats consist of large pores (>30 μm) that are formed and primarily occupied by roots or detritus, aka biopores, leading to high supply of C and nutrients, and O_2 availability with temporarily variable moisture conditions. The bacterial communities in Lp are dominated by plant-residue decomposers and abundant predators, and have low potential for C sequestration because of intensive decomposition of organics. (2) The SpRich habitats consist of small pores (<10 μm) in rhizosphere and detritusphere, thus with high supply of C and nutrients but somewhat limited O_2 and with a highly diverse bacterial community dominated by microbial-residue decomposers. The SpRich habitats have high potential for C sequestration because of high C input, fast microbial turnover, and close contact of microbial necromass with soil matrix. (3) The SpPoor habitats consist of small pores (<10 μm) in the bulk soil (far from roots or detritus), with limited C and nutrient resources and low O_2 leading to the bacterial community of oligotrophic passive consumers. The SpPoor habitats have low potential for C sequestration because of low C input and slow microbial turnover. Detailed description of the hypothesized characteristics of the three habitats are provided in Table 1.

also get occasional O_2 influxes. The SpPoor habitats are located in the bulk soil, getting only occasional new C and nutrient inputs with inflowing soil solution, which promotes frugal C and N use strategies by bacterial inhabitants.

We postulate that the three microhabitats differ in ecological C-acquisition strategies of their resident microorganisms. The four proposed C-acquisition groups⁶ are (i) plant polymer decomposers (1^o decomposers), (ii) microbial necromass decomposers (2^o decomposers), (iii) predators that consume living microorganisms (predatory microbes), and (iv) passive consumers assimilating DOC, representing a spectrum of life strategies from oligotrophic to copiotrophic⁶. Our findings suggest that:

- (1) Plant polymer decomposers outnumber microbial necromass decomposers in the Lp habitats, while the converse is true for SpRich habitats. Several of the taxa found exclusively in the large pores (Supplementary Table S7), i.e. *Celvibrio*, *Chitinophaga*, and *Sphingobium*, are known to be capable of plant-residue consumption via cellulolysis, lignolysis, and chitinolysis (FAPROTAX⁶⁸).
- (2) Predators are more abundant in Lp than in SpRich habitats and are very rare or absent in SpPoor⁶⁹, (refs in ref. 67). *Bacteriovorax* was ¹³C-enriched exclusively in the large pores; and multiple members of *Bdellovibrio* taxon were among top-enriched organisms in the large pores, with only one low enriched *Bdellovibrio* OTU found in the small pores of the prairie system (Supplementary Table S6).

- (3) While passive DOC consumers are present in all soil habitats, in Lp and SpRich they are fast-growing copiotrophs quickly responding to frequent inputs of labile substrates, whereas slow-growing oligotrophic microbes^{70,71}, adapted to low inputs, dominate the SpPoor habitat. For example, *Pseudomonas*, *Burkholderiaceae*, and *Pseudarthrobacter* are putative copiotrophs and *r*-strategists^{71,72}. These organisms strongly and instantly responded to glucose addition to both large and small pores (Fig. 3b and Supplementary Table S5). Even 30 days later they remained among the abundant enriched groups in the large pores of either both or at least one of the plant systems.

Micro-environmental differences and specifics in C-acquisition strategies imply distinctive functional roles for the three microhabitats in soil C processing and protection (Table 1). Lp and SpRich habitats are locations where microbial communities are ready for active processing of frequent new C inputs. SpRich and SpPoor habitats with their large contact areas between organic compounds and mineral surfaces provide C protection. However, C protection in pores of the SpPoor habitat can be facilitated even further due to the SpPoor habitat's lower C saturation and the relative inability of its microbial inhabitants to profit from the ¹³C consumed by the so-called first-responder community⁷³—the phenomenon reflected in this study's particularly high ¹³C in the small pore of the switchgrass soil (Fig. 2b). Yet, in natural conditions the inauspicious pore architecture of SpPoor habitats and their low pore connectivity limits delivery of new C. Thus, the high C protection potential of SpPoor habitats does not translate into tangible soil C gains.

The proposed *micro-habitat concept* described here is a first step towards a generalizable C processing classification of conditions within an intact, undisturbed soil matrix. While we recognize that process-based modeling of soil C cycling requires capturing the contributions of key, highly variable microbial drivers^{5,65,74}, thus far only a crude rhizosphere vs. bulk soil classification of such conditions has been used in modeling⁷⁵. Our work provides experimental evidence linking physical and biochemical properties of the soil matrix with microbial functional traits and C transformations in soil microhabitats in situ. Moreover, the specific locations and sizes of the microhabitats can be quantified using X-ray μCT -based POM and pore data, thus enabling quantification and modeling of the unique physical, biochemical, biological, and ecological contributions of each habitat towards key soil functions.

Future work is needed to further test this concept. In particular there is a need to quantify microhabitats across a wide range of soil types with various textural and mineralogical characteristics, expanding to soils under specific plant communities and management practices. That said, results here suggest that this model should be fairly robust, and could represent a powerful additional way to characterize soil microhabitats in a way that also explains the distribution of microbial taxa and their important ecosystem-level processes, leading to improved biogeochemical models and perhaps management interventions.

Methods

Field experiment and soil sampling

Soil samples were collected from the Cellulosic Biofuel Diversity Experiment located at the Kellogg Biological Station Long-term Ecological Research Site⁷⁶, Hickory Corners, Michigan. The studied soil is a mesic Typic Hapludalf with 52% sand, 39% silt, and 9% clay⁷⁷. The experiment is a randomized complete block (RCB) design established in 2008. For this study, we selected three plant communities that represent a range of diversities: a multiyear fallow system where soil was kept free of vegetation during 2008–2016 and then planted to corn 2016–2018; a monoculture switchgrass (variety “Southlow”) system;

Table 1 | Hypothesized physical, biochemical, and microbial community characteristics of the identified microhabitats: large pore (Lp), substrate-rich small pores (SpRich), and substrate-poor small pores (SpPoor)

Micro-habitat characteristics		Micro-habitat type
		Large pores (30–150 µm Ø)
Physical properties	O ₂ availability	- Oxic conditions prevail most of the time, with anoxic micro-sites in areas of active decomposition.
	Water availability	- Water availability is highly spatially and temporally variable: ranges from water bodies with high spatial continuity enabling microbial transport when wet to isolated menisci of varying thickness when dry; subject to frequent wet/dry cycles.
Spatial connectivity		- Pores are well-connected. - Water bodies within the pores are well-connected.
Biochemical properties	C and N inputs as particulate and dissolved matter	- High and frequent C inputs from plant roots, detritus, and DOC from inflowing water. - High N availability from plant inputs and decomposing residues.
Bacterial communities	Diversity	Low
	O ₂ status	- Strict and facultative aerobes, few to no strict anaerobes
	C-acquisition strategy	- Plant-residue decomposers are more abundant than microbial-residue decomposers. - Predators are abundant. - Passive consumers are fast-growing copiotrophs.
Metabolic pathways	Glycolysis/citric acid cycle	- Citric acid cycle. - Fatty acid synthesis.
	Nitrogen	- Glutamate synthesis. - When anaerobic, denitrification dominates N processing.
	Methane	
Implications for soil C storage		- New C inputs are quickly processed, and large quantities of the new C continue to be utilized by the microbial community. - The DOC products have a high potential to be carried out of the habitat by convective transport. - Some DOC products may diffuse into surrounding soil matrix, to be protected from further decomposition. - C processing as precursor to storage. - Organic C accrual is mainly by partly processed POM, which is stored for short- and medium-term
		- New C inputs are quickly processed, and large quantities of the new C continue to be utilized by the microbial community. - Large quantities of DOC can diffuse into surrounding soil matrix, to be protected from further decomposition. - Organic C accrual is mainly by microbial necromass, which is stored for long periods. - C storage is maximal.
		- New C inputs are not as efficiently and completely utilized by the microbial community as in the other habitats. - Unclaimed DOC can diffuse into surrounding soil matrix, to be protected from further decomposition. - Organic C accrual is solely by microbial necromass, which is stored for very long periods. - C storage potential is high, but the actual C storage is low.
		- Pores are poorly connected. - Water bodies are spatially fragmented.
		- Low and infrequent C inputs as DOC brought by convective water influx. - Low N availability.
		- Pores are partially connected. - Water bodies are spatially fragmented.
		- Medium quantities but high frequencies of C inputs from root hairs and DOC diffusing from roots and decomposing detritus as well as brought by convective water influx. - Medium N availability from plant inputs and decomposing residues.
		High
		- Facultative aerobes, strict anaerobes
		- Microbial-residue decomposers are present. - Plant-residue decomposers and predators are absent. - Passive consumers are slow-growing oligotrophs.
		- Wide use of degradation products. - Glycolysis. - Protein maintenance.
		- When anaerobic, dissimilatory nitrate reduction dominates N processing. - N fixation is used to counteract N shortages.
		Prevalent.

Also presented are the prevalence of microorganisms with specific carbon acquisition ecological strategies⁸ within each habitat, and hypothesized contributions of each habitat to soil C cycling.

and a high plant diversity system consisting of 6 native grasses and 24 native forbs typical of Michigan restored prairie communities. We refer to these systems as bare, switchgrass, and prairie, respectively. The RCB design enables us to establish causality for effects of different systems on soil and microbial characteristics insofar as the three systems were randomly assigned to 9×27 m plots within replicate blocks. The entire experimental area was located on a well-drained flat terrain, in a field cultivated conventionally for >100 years previous.

Three experimental plots per switchgrass and prairie vegetation system and four plots per bare system were sampled in 2019. In each plot six intact soil cores (5 cm \varnothing , 2.5 cm height) were collected from the 5 to 7.5 cm depth increment using a 5.7 cm diameter soil core sampler (Soilmoisture Equipment Corp., Santa Barbara, CA, USA) fitted with an acrylic sleeve. After removing the 0–5 cm top soil layer the sampler was gradually driven into the soil to collect a minimally disturbed intact core. The intact cores collected from each plot were in close proximity (<3 cm distance) to each other. The loose (bulk) soil around the cores was also collected for subsequent analyses (~500 g per plot). To prevent drying the cores were wrapped in parafilm and aluminum foil and the bulk soil was placed into zip-lock bags, all stored at 4 °C prior to analyses. Gravimetric water content was measured using a 20 g sub-sample of the bulk soil immediately upon collection. A workflow diagram is provided in Supplementary Fig. S1a.

Bulk soil analyses included total C and N via combustion analysis⁷⁸, soil pH⁷⁹, available phosphorus and potassium⁸⁰, and cation exchange capacity⁸¹. Root biomass in each plot was measured using the core method⁸², and bulk density⁸³ was obtained from the spare (6th) intact core of each plot (Supplementary Fig. S1a).

Glucose addition experiment

The ¹²C- and ¹³C-glucose addition experiment was built on the matrix potential approach of substrate additions to the soil pores of contrasting sizes^{42–44}. We used three glucose addition treatments: (1) glucose dissolved in DI water added to the 4–10 μm \varnothing pores of the intact soil cores (referred to as the small-pore treatment), (2) dissolved glucose added to the 30–150 μm \varnothing pores (referred to as the large-pore treatment), and (3) a control treatment with only water added to corresponding pore sizes. We used a set of 5, out of the 6, cores collected from each experimental field plot, with one core assigned to the control treatment without glucose application; two cores assigned to the small and the other two cores assigned to the large-pore treatments, with either ¹²C or ¹³C-glucose added to each core (Supplementary Fig. S1a).

For glucose additions five cores representing a full set of glucose treatments from each experimental plot were processed simultaneously (Supplementary Fig. S1b). We used 100 mg ml⁻¹ solution of either ¹²C glucose or ¹³C-glucose (99 atom %). Equal volume (0.8 ml) of the solution was applied to every core, resulting in 80 mg of glucose added per core. Thus, we achieved an application rate of roughly 50 $\mu\text{mole C g}^{-1}$ dry soil, consistent with reported recommendations for SIP analyses⁸⁴. It should be noted that given high variability in pore space, stone contents, root and other organic residues contents within the intact cores of this study, adding equal amounts of glucose per soil volume, as opposed to per soil mass, was regarded as the only feasible approach that would enable consistent comparisons among the studied treatments. The volume of the solution to be added (0.8 ml) was estimated as the average smallest volume of the pores of the target group, which happens to be the small pores in the switchgrass cores (Supplementary Table S2), as observed from μCT analysis. Adding this much liquid would, in theory, fill completely all target small pores in those soil cores where the volume of such pores was the smallest, while the cores with greater volumes of small pores would have some of those target pores left without the glucose solution. Likewise, since large pores occupied a much bigger soil volume than the small pores (Supplementary Table S2), only a portion of such pores has received

the glucose solution. Thus, the differences in the actual volumes of small or large pores among the studied treatments were irrelevant for the treatment comparisons—the same volumes of pores of each size group were filled with the same glucose solution in each core. It is important to point out that the only possible effect of the discrepancies between the volume of the added glucose solution and the actual volumes of the target pores in this study would be a decrease and some dampening of the differences between the pore treatments and the systems. Thus, our results, as far as comparisons between the pore treatments and systems, represent conservative estimates of the real differences that could have been obtained, if it were possible to perfectly fill all pores of the target size ranges.

To ensure glucose delivery to the pores of specific size ranges, we used the following approach (Supplementary Table S1): first, cores were drained to 400 kPa to ensure that all pores >1 μm were emptied, then water was added to bring them to 75 kPa to fill the smallest (1–4 μm) pores with water. It should be noted that while draining the soil always results in some rearrangement of soil pore space, the relatively coarse texture and non-expanding clay mineralogy of the studied soil minimized such impact. Then either water (in the large-pore and control treatments) or glucose solution (in the small-pore treatment) was added to bring the samples to 30 kPa, filling the ~4–10 μm pores. Then water was added to all cores to 10 kPa in order to create a water buffer between the target small and large pores. Finally, either water (in the small-pore and control treatments) or glucose (in the large-pore treatment) was added to 2 kPa to fill large (~30–150 μm) pores. Glucose additions were conducted in a cold room (2 °C) to minimize microbial consumption of the added glucose until it reached the target pores, after which the cores were brought to 22 °C for incubation.

To assess the microbial community involved in the immediate consumption of the added glucose and then to explore the longer-term fate of the added C, we incubated the soil cores for either 24 h or 30 days (Supplementary Fig. S1a). Each incubation period was followed by analysis of $\delta^{13}\text{C}$ in the soil, dissolved organic carbon (DOC), and microbial cytosol, as well as in gas samples taken during the subsequent 30-day incubation. For the first incubation the entire cores were placed into incubation jars and kept there for 24 h at 22 °C. Then the cores were cut in half, one half was passed through a 2-mm sieve, with roots and stones removed, and frozen at -20 °C for subsequent SIP and microbial biomass C analyses, with part of the soil retained to determine soil ¹³C (Supplementary Fig. S1a). The other half was returned to the incubation jar and incubated for 29 more days, after which it was also sieved and frozen. Gas samples for ¹³CO₂ and total CO₂ analyses were taken from the jars at 1, 3, 7, and 30 days during the incubation. Each incubation jar was equipped with a small water-holding container to maximize air humidity and minimize evaporation. Soil gravimetric moisture content was measured using a sub-sample after both 24 h and 30-day incubations. The average moisture contents after the 24 h and 30-day incubations were equal to 24% and 22%, respectively.

Microbial biomass was analyzed using the fumigation-extraction method^{85,86}. In brief, soil samples fumigated in ethanol-free chloroform and non-fumigated samples (5 g) were subjected to 0.5 M KCl extraction (1:5 of soil:solution ratio). Upon shaking, centrifugation, and filtering (0.45 μm membrane), the extracts were freeze-dried and subjected to the total C and $\delta^{13}\text{C}$ analyses. The difference between fumigated and non-fumigated samples was reported as cytosol C, while the non-fumigated samples represented dissolved organic C. Measurements of $\delta^{13}\text{C}$ were performed in the stable isotope facility at Michigan State University using an Isoprime Vision IRMS interfaced to a Vario Isotope Cube elemental analyzer (Elementar).

X-ray computed microtomography (μCT) scanning

One of the six cores collected at each experimental plot was used to characterize soil pores and to quantify particulate organic matter

(POM) within the intact soil matrix using X-ray μ CT. Two intact mini-cores (8 mm \varnothing and 1 cm height) were taken from each core (Supplementary Fig. S1a), air-dried, and scanned at the Advanced Photon Source, Argonne National Laboratory (sector 13-BM-D). The scanning energy of the monochromatic beam was 24 keV, and the scanning resolution was $\sim 4 \mu\text{m}$. The μ CT image analyses were conducted using ImageJ/Fiji^{87,88}. The noise on images was removed using a 3D median filter, followed by contrast enhancement with 0.3 saturation.

Soil POM was determined using a volumetric approach of X-ray μ CT, which has demonstrated good agreement with conventional POM measurements^{51,89,90}. POM fragments were identified using the machine-learning-based classification approach in ilastik 1.0⁹¹ (Fig. 1a); then the volumes of all identified POM fragments were added and the overall POM was reported as a percent of the total soil volume. Size distributions of POM fragments were obtained using the particle analyzer tool of BoneJ⁹². The smallest size of the POM fragments was set as equal to $\times 10$ of the image resolution, e.g., $40 \times 40 \times 40 \mu\text{m}$, as a conservative estimate of the size of the reliably detectable μ CT image features^{93,94}. The identified POM was regarded as solid material in subsequent solid-pore segmentations.

The solids and pores in the images were segmented using the Otsu method as implemented in ImageJ⁸⁸. Pore-size distributions were obtained using the maximally inscribable spheres approach as implemented in Xlib plug-in for ImageJ⁹⁵. The connectivity of the pore space was determined as the percent of the total soil volume that was occupied by the largest cluster of interconnected pore voxels⁹⁶. Then we also determined the volumes occupied by clusters of small ($4\text{--}10 \mu\text{m} \varnothing$) and large ($30\text{--}150 \mu\text{m} \varnothing$) pores not connected with the largest pore cluster.

We used the average distances to pores and POM, i.e., the average distances between individual voxels of the soil solid material and the border of the nearest visible ($\varnothing > 10 \mu\text{m}$) pore or of the nearest POM fragment, as a measure of homogeneity in pore and/or POM spatial distributions. The average distances between solid soil matrix voxels and either pores or POM were obtained using a 3D distance function.

We also determine the percent of the soil matrix volume that was within the active detritusphere, i.e., in close proximity to POM fragments. While the overall distance at which roots or decomposing plant residues can influence the surrounding soil matrix can extend to $3\text{--}4 \text{mm}$ ^{97–99}, here we selected $< 300 \mu\text{m}$ as the distance where detritus influences and microbial responses to detritus were expected to be strongest^{97,100}. While in the scanned air-dried samples there were often pockets of empty pore space around the POM fragments, we assumed that under natural conditions of a relatively wet soil the POM fragments would act as sponges^{101–103} expanding their volumes and filling most of the voids they are in, while the remaining empty areas around the fragments would be occupied by water films. We build on this assumption while estimating the extent of the influence of POM on the surrounding soil, which is assumed to proceed via both direct contact with minerals and through the water films.

We employed multi-energy X-ray computed tomography (μ CT) scanning to visualize the specific locations of and volumes occupied by the liquids added to pores of different sizes. Specifically, to characterize the distributions of glucose solutions when added to the cores, three intact mini-cores (8 mm \varnothing and 5 mm height) were subjected to multi-energy X-ray μ CT scanning after adding the dopant solutions (10% KI and 10% BaCl_2) to the small and large soil pores¹⁰⁴. Since glucose at the concentrations used in this study does not precipitate and the water-glucose solution has a similar viscosity to the KI and BaCl_2 solutions, we assume that the distribution of glucose in soil pores does not differ significantly from those for the dopant solutions when applied in the same volumes and at the same matrix potentials. The spatial distribution of the dopants in soil pores was quantified based on the dopant mass attenuation coefficients. Even though the time-consuming and expensive nature of the analysis limited the

number of cores could be analyzed, we were able to verify where the applied solutions were placed within the soil pore system when applied to target small and large pores, as well as to characterize the hydraulic connectivity of large and small pores (Fig. 1b and Supplementary Fig. S2). Connectivity of the pore space filled with the solutions targeting small or large pores was estimated as the size of the largest cluster 90% saturated by Iodine for small pores or as the size of the largest cluster 90% saturated by Ba for large pores. This approach was proposed and first implemented in the Scamp plug-in for ImageJ¹⁰⁵. Here we used Particle Analyzer function of the BoneJ plug-in to identify clusters and determine their sizes.

DNA extraction, ultracentrifugation, and fractioning

DNA was extracted using the DNeasy PowerSoil kit or the DNeasy PowerSoil Pro kit (Qiagen, USA) following the manufacturer's protocols. Approximately 1 g of each soil was used for DNA extraction. Carnation instant nonfat dry milk (40 mg, Nestlé, Rosslyn, VA) was added at the beginning of the extraction process to improve the DNA yield¹⁰⁶. DNA concentrations from each extraction were determined using the Qubit fluorometer (Thermo Fisher, USA) with the dsDNA HS Assay kit. For ultracentrifugation, approximately $10 \mu\text{g}$ of each DNA extract was mixed with Tris-EDTA buffer (10 mM Tris, 1 mM EDTA, pH 8) and cesium chloride (CsCl) solution (1.62 M) and loaded into Quick-Seal Round-Top Polypropylene tubes ($13 \times 51 \text{mm}$, 5 ml; Beckman Coulter, USA). Refractive index (RI) values of each solution were determined using AR200 digital refractometer (Leica Microsystems Inc., Buffalo Grove, IL) and the RI was adjusted to between (1.4069–1.4071) by adding small volumes of TE buffer or CsCl solution. The sealed tubes were ultracentrifuged at $178,000 \times g$ (20°C) for 46 h in a StepSaver 70 V6 vertical titanium rotor (8 by 5.1 mL capacity) within a Sorvall WX 80 Ultra Series centrifuge (Thermo-Scientific, Waltham, MA). Following ultracentrifugation, each tube was placed onto a fraction collection system (Beckman Coulter) to generate ~ 26 fractions ($200 \mu\text{L}$). The RI of each fraction was determined, and, from this, buoyant density values were calculated. CsCl in the fractions was removed using linear polyacrylamide (Thermo-Scientific, USA) and a polyethylene glycol solution (1.6 M NaCl, 30% PEG solution; Thermo-Scientific, USA). The DNA concentration in each fraction was determined using the dsDNA HS Assay kit to identify the four heaviest fractions with the minimum amount of DNA for high throughput sequencing. For each of the labeled and unlabeled glucose-amended samples, sixteen tubes were ultracentrifuged: four replicate blocks for the bare soil 24 h incubation; three replicate blocks for both the switchgrass 24 hour incubation and the high diversity prairie 24 h incubations; and three replicate blocks for both the switchgrass 30-day incubation and the high diversity prairie 30-day incubations. As both small pores and large-pore incubations were also examined, in total, 64 tubes were ultracentrifuged (2 glucose forms [^{12}C and ^{13}C] $\times 16$ treatments/blocks $\times 2$ pore sizes).

Miseq Illumina sequencing and Mothur analysis

Total genomic DNA extracts (before ultracentrifugation) and ultracentrifugation fractions were submitted to the Research Technology Support Facility (RTSF) at MSU for 16 S rRNA gene amplicon sequencing. For each of the 64 ultracentrifugation runs (as described above), three heavy fractions (buoyant density 1.73–1.75 g/ml) and one light fraction ($\sim 1.70 \text{g/ml}$) were submitted in triplicate for sequencing. This involved amplification of the V4 region of the 16 S rRNA gene using dual indexed Illumina compatible primers 515f/806r, as previously described¹⁰⁷. PCR products were batch normalized using Invitrogen SequalPrep DNA Normalization plates and the products recovered from the plates pooled. The pool was cleaned and concentrated using AmpureXP magnetic beads; then QC'd and quantified using a combination of Qubit dsDNA HS, Agilent 4200 TapeStation HS DNA1000, and Kapa Illumina Library Quantification qPCR assays. The pool was

loaded onto an Illumina MiSeq v2 standard flow cell and sequencing was performed in a 2 × 250 bp paired end format using a MiSeq v2 500 cycle reagent cartridge. Custom sequencing and index primers were added to appropriate wells of the reagent cartridge.

Base calling was performed by Illumina Real Time Analysis (RTA) v1.18.54 and RTA output demultiplexed and converted to FastQ format with Illumina Bcl2fastq v2.19.1. The amplicon sequencing data in the fastq format was analyzed by Mothur¹⁰⁸ using the Mothur MiSeq SOP (accessed August 2021)¹⁰⁷. Briefly, the Mothur analysis involved trimming the raw sequences and quality control. The SILVA bacteria database (Release 138) for the V4 region¹⁰⁹ was used for the alignment. Chimeras, mitochondrial, and chloroplast lineage sequences were removed, then the sequences were classified into operational taxonomic units (OTUs) at a 0.03 cutoff. The OTUs were then grouped into taxonomic levels and the downstream analysis conducted in R (version 4.0.2)¹¹⁰ with RStudio (version 1.5042)¹¹¹. The sequencing data of the total DNA and SIP fractions were submitted to NCBI under Bioproject PRJNA801760 (accession numbers SAMN25378717 to SAMN25378796) and Bioproject PRJNA802612 (accession numbers SAMN25563888 to SAMN25564655), respectively. Sequencing data from the total DNA extracts and the ultracentrifugation fractions were analyzed separately, as described below.

Total DNA community analysis

For the analysis of the total DNA samples, two Mothur files (shared file and taxonomy file) along with an independently created metafile were used as the input for packages phyloseq¹¹² (version 1.34.0), file2meeco (version 0.1.0)¹¹³, and microeco (version 0.5.1)¹¹³. A total of 28.1 Gbytes were sequenced, with an average of 31.5 Mbytes for each sample. This resulted in the creation of (1) a phylum-level bar chart, (2) a Venn diagram of OTU abundance, and (3) two boxplots at the genus and order level. The packages phyloseq¹¹² (version 1.34.0), microbiome¹¹⁴ (version 1.12.0) and ampvis2 (version 2.7.11)¹¹⁵ were used to (1) generate heatmaps of the most abundant genera, (2) perform alpha diversity analysis (Chao1, ACE, Shannon's values, Simpson, Inverse Simpson, and Fisher indices), and (3) create barplots for the most abundant classes. The data for the alpha diversity measurements were rarefied using the phyloseq function `rarefy_even_depth(pseq, sample.size = 50, rngseed = 1)`. The "adonis" function in the package vegan (version 2.5.7)¹¹⁶ was used to test differences between microbial communities in different soil treatments with Permutational Multivariate Analysis of Variance (PERMANOVA). The "pairwise.adonis" function in package pairwiseAdonis (version 0.4)¹¹⁷ was used for the comparison of significant PERMANOVA results ($P < 0.05$). The "simper" function in the package vegan (version 2.5.7)¹¹⁶ was used for dissimilarity analyses for the significant comparison results ($P < 0.05$). The abundance and the classification of top twenty OTUs contributing to the difference between treatments and the connection between the OTUs and different samples was determined using the circlize package (version 0.4.13)¹¹⁸.

Identification of enriched phylotypes

Sequencing datasets were compared between the heavy and light fractions of the ¹³C-glucose-amended samples and fractions of similar buoyant density from the ¹²C glucose-amended samples to determine which phylotypes were responsible for label uptake. For this, data generated from the packages phyloseq¹¹² (version 1.34.0) and microbiome¹¹⁴ (version 1.12.0) were analyzed using the packages dplyr (version 1.0.7)¹¹⁹, tidyr (version 1.1.4)¹²⁰, ggpubr (version 0.4.0)¹²¹ and rstatix (version 0.7.0)¹²². Specifically, those enriched in the heavy fractions of the ¹³C-glucose-amended samples (compared to the same fractions in the ¹²C glucose-amended samples) were determined using the Wilcoxon test (function `wilcox_test`) in RStudio (one-sided, $P < 0.05$). From those significantly enriched, the six most abundant

were selected for the creation of boxplots using ggplot2 (version 3.3.5)¹²³. The analysis also included the comparison of phylotypes in the light fractions of the ¹³C-glucose-amended fractions compared to the light fractions of the ¹²C glucose-amended samples. Those enriched in the light fractions of the ¹³C-glucose-amended samples were removed from the above analysis to limit the possibility of reporting false positives.

Function prediction by PICRUSt2

PICRUSt2¹²⁴ was used to predict the microbial functions of the sequencing data from the Kyoto Encyclopedia of Genes and Genomes (KEGG) orthologs (KO)¹²⁵. Biom and fasta files generated by Mothur were used for this analysis. The PICRUSt2 analysis included sequence placement with EPA-NG¹²⁶ and gappa¹²⁷, hidden state prediction with castor R package¹²⁸ and pathway abundance inference with MinPath¹²⁹.

The KO functions associated with carbohydrate metabolism and energy metabolism were examined, including citrate cycle (TCA cycle) [PATHko00020], glycolysis gluconeogenesis [PATHko00010], methane metabolism [PATHko00680] and nitrogen metabolism [PATHko00910]. The metagenome output files were analyzed with ggplot2 (version 3.3.5)¹²³ and ggpubr (version 0.4.0)¹²¹ in RStudio (version 1.5042)¹¹¹. The relative abundance of genes associated with carbohydrate metabolism and energy metabolism were also determined for each treatment. The enriched genes in the samples amended with ¹³C-glucose were determined using a similar approach as described above for the enriched phylotypes (Wilcoxon test, $p < 0.05$). The eight most abundant genes associated with carbohydrate metabolism and energy metabolism were displayed in barplots. The most abundant phylotypes associated with methane metabolism and nitrogen metabolism for the total DNA samples were determined in RStudio (version 1.5042)¹¹¹.

Statistical analysis for soil and CO₂ emission data

Statistical analyses were conducted using the mixed model approach implemented in *proc mixed* procedure of SAS¹³⁰. The statistical models for all variables included the fixed effect of the plant system and pore treatment, with their interaction, and the random effect of the field replication blocks. The normality of residuals and equal variance assumptions were checked by examining normal probability plots and side-by-side boxplots, respectively. When normality assumption was found to be violated the data were log-transformed. When the equal variance assumption was found to be violated an unequal variance analysis was conducted using *repeated* option of *proc mixed*. When the main effects were statistically significant ($P < 0.05$) mean separations among the systems were conducted using *t* tests, while statistically significant interactions were followed by slicing. Differences significant at $P < 0.1$ are reported as trends. Detailed descriptions of the specific statistical models used for the studied variables along with their ANOVA results are provided in respective supplementary tables (Supplementary Tables S3 and S4).

The statistical models for the CO₂ emission data and other data from the incubation experiments included the incubation time and its interactions with plant systems and pore treatments, as fixed effects. The models for these data and for the data from μ CT image analyses also included the random effects of experimental plots (nested within the plant systems). Time was treated as a repeated measures factor in the analyses for the ¹³CO₂ and CO₂ data, and appropriated repeated measures structures were selected as per Milliken and Johnson¹³⁰ (via Bayesian Information Criterion) and used for subsequent plant system and pore treatment comparisons.

Reporting summary

Further information on research design is available in the Nature Portfolio Reporting Summary linked to this article.

Data availability

The sequencing data of the total DNA and SIP fractions were submitted to NCBI under Bioproject PRJNA801760 (accession numbers SAMN25378717 to SAMN25378796) and Bioproject PRJNA802612 (accession numbers SAMN25563888 to SAMN25564655), respectively. The sequences were aligned using SILVA bacteria database (Release 138) obtained from the Miseq SOP (accessed August 2021). All the data is publicly released. Source data are provided with this paper.

References

- Lal, R., Monger, C., Nave, L. & Smith, P. The role of soil in regulation of climate. *Philos. Trans. Royal Soc. B* **376**, 20210084 (2021).
- Rillig, M. C. et al. The role of multiple global change factors in driving soil functions and microbial biodiversity. *Science* **366**, 886–890 (2019).
- Gentry, T. J., Fuhrmann, J. J. & Zuberer, D. A. *Principles and Applications of Soil Microbiology*, 3rd edn. (Elsevier, 2021).
- Vos, M., Wolf, A. B., Jennings, S. J. & Kowalchuk, G. A. Micro-scale determinants of bacterial diversity in soil. *FEMS Microbiol. Rev.* **37**, 936–954 (2013).
- Nunan, N., Schmidt, H. & Raynaud, X. The ecology of heterogeneity: soil bacterial communities and C dynamics. *Philos. Trans. Royal Soc. B Biol. Sci.* **375**, 20190249 (2020).
- Morrissey, E. M. et al. Carbon acquisition ecological strategies to connect soil microbial biodiversity and carbon cycling. *Soil Biol. Biochem.* **177**, 108893 (2023).
- Morris, S. J. & Blackwood, C. B. 8—THE ECOLOGY OF SOIL ORGANISMS. In *Soil Microbiology, Ecology and Biochemistry*, third edition (ed. Paul, E. A.) 195–230 (Academic Press, 2007).
- Panikov, N. S. *Microbial Growth Kinetics* (Chapman & Hall, 1995).
- Koebnick, N. et al. High-resolution synchrotron imaging shows that root hairs influence rhizosphere soil structure formation. *N. Phytol.* **216**, 124–135 (2017).
- Birt, H. W. G., Tharp, C. L., Custer, G. F. & Dini-Andreote, F. Root phenotypes as modulators of microbial microhabitats. *Front. Plant Sci.* **13**, 1003868 (2022).
- Zhang, N. et al. Theory of microbial coexistence in promoting soil–plant ecosystem health. *Biol. Fert. Soils* **57**, 897–911 (2021).
- Lucas, M., Schlüter, S., Vogel, H.-J. & Vetterlein, D. Roots compact the surrounding soil depending on the structures they encounter. *Sci. Rep.* **9**, 16236 (2019).
- Carson, J. K. et al. Low pore connectivity increases bacterial diversity in soil. *Appl. Environ. Micro.* **76**, 3936–3942 (2010).
- Chenu, C., Hassink, J. & Bloem, J. Short-term changes in the spatial distribution of microorganisms in soil aggregates as affected by glucose addition. *Biol. Fert. Soils* **34**, 349–356 (2001).
- Sleutel, S. et al. Manipulation of the soil pore and microbial community structure in soil mesocosm incubation studies. *Soil Biol. Biochem.* **45**, 40–48 (2012).
- Ravenek, J. M. et al. Long-term study of root biomass in a biodiversity experiment reveals shifts in diversity effects over time. *Oikos* **123**, 1528–1536 (2014).
- Eisenhauer, N. et al. Plant diversity effects on soil microorganisms support the singular hypothesis. *Ecology* **91**, 485–496 (2010).
- Zak, D. R., Holmes, W. E., White, D. C., Peacock, A. D. & Tilman, D. Plant diversity, soil microbial communities, and ecosystem function: are there any links? *Ecology* **84**, 2042–2050 (2003).
- Lange, M. et al. Plant diversity increases soil microbial activity and soil carbon storage. *Nat. Commun.* **6**, 6707 (2015).
- Kravchenko, A. N. et al. Microbial spatial footprint as a driver of soil carbon stabilization. *Nat. Commun.* **10**, 3121(2019).
- Borer, B., Tecon, R. & Or, D. Spatial organization of bacterial populations in response to oxygen and carbon counter-gradients in pore networks. *Nat. Commun.* **9**, 769 (2018).
- Schluter, S. et al. Denitrification in soil aggregate analogues—effect of aggregate size and oxygen diffusion. *Front. Environ. Sci.* **6**, 17 (2018).
- Monga, O., Bousso, M., Garnier, P. & Pot, V. 3D geometric structures and biological activity: application to microbial soil organic matter decomposition in pore space. *Ecol. Model.* **216**, 291–302 (2008).
- Kravchenko, A. N. et al. Intra-aggregate pore structure influences phylogenetic composition of bacterial community in macro-aggregates. *Soil Sci. Soc. Am. J.* **78**, 1924–1939 (2014).
- Xia, Q., Zheng, N. G., Heitman, J. L. & Shi, W. Soil pore size distribution shaped not only compositions but also networks of the soil microbial community. *Appl. Soil Ecol.* **170**, 104273 (2022).
- Benucci, G. M. N. et al. The microbiome structure of decomposing plant leaves in soil depends on plant species, soil pore sizes, and soil moisture content. *Front. Microbiol.* **14**, 1172862 (2023).
- Tecon, R. & Or, D. Biophysical processes supporting the diversity of microbial life in soil. *FEMS Microbiol.* **41**, 599–623 (2017).
- Long, T. & Or, D. Aquatic habitats and diffusion constraints affecting microbial coexistence in unsaturated porous media. *Water Resour. Res.* **41**, (2005).
- Long, T. & Or, D. Microbial growth on partially saturated rough surfaces: simulations in idealized roughness networks. *Water Resour. Res.* **43**, (2007).
- Long, T. & Or, D. Dynamics of microbial growth and coexistence on variably saturated rough surfaces. *Micro. Ecol.* **58**, 262–275 (2009).
- Or, D., Smets, B. F., Wraith, J. M., Dechesne, A. & Friedman, S. P. Physical constraints affecting bacterial habitats and activity in unsaturated porous media—a review. *Adv. Water Resour.* **30**, 1505–1527 (2007).
- Bailey, V. L., Smith, A. P., Tfaily, M., Fanster, S. J. & Bond-Lamberty, B. Differences in soluble organic carbon chemistry in pore waters sampled from different pore size domains. *Soil Biol. Biochem.* **107**, 133–143 (2017).
- Bouckaert, L. et al. Carbon mineralisation and pore size classes in undisturbed soil cores. *Soil Res.* **51**, 14–22 (2013).
- Smith, A. P. et al. Shifts in pore connectivity from precipitation versus groundwater rewetting increases soil carbon loss after drought. *Nat. Commun.* **8**, 1335 (2017).
- Bailey, V. L., Pries, C. H. & Lajtha, K. What do we know about soil carbon destabilization? *Environ. Res. Lett.* **14**, 083004 (2019).
- Lucas, M., Vetterlein, D., Vogel, H.-J. & Schlüter, S. Revealing pore connectivity across scales and resolutions with X-ray CT. *Eur. J. Soil Sci.* **72**, 546–560 (2021).
- Schlüter, S., Eickhorst, T. & Mueller, C. W. Correlative imaging reveals holistic view of soil microenvironments. *Environ. Sci. Technol.* **53**, 829–837 (2019).
- Anderson, J. P. E. & Domsch, K. H. A physiological method for the quantitative measurement of microbial biomass in soils. *Soil Biol. Biochem.* **10**, 215–221 (1978).
- Lemanski, K. & Scheu, S. Incorporation of ¹³C labelled glucose into soil microorganisms of grassland: effects of fertilizer addition and plant functional group composition. *Soil Biol. Biochem.* **69**, 38–45 (2014).
- Kong, Y. et al. Are the microbial communities involved in glucose assimilation in paddy soils treated with different fertilization regimes for three years similar? *J. Soil Sediment.* **18**, 2476–2490 (2018).
- Padmanabhan, P. et al. Respiration of ¹³C-labeled substrates added to soil in the field and subsequent ¹⁶S rRNA gene analysis of ¹³C-labeled soil DNA. *Appl Environ. Micro.* **69**, 1614–1622 (2003).

42. Killham, K., Amato, M. & Ladd, J. N. Effect of substrate location in soil and soil pore-water regime on carbon turnover. *Soil Biol. Biochem.* **25**, 57–62 (1993).
43. Kravchenko, A., Guber, A., Gunina, A., Dippold, M. & Kuzyakov, Y. Pore-scale view of microbial turnover: combining ^{14}C imaging, μCT and zymography after adding soluble carbon to soil pores of specific sizes. *Eur. J. Soil Sci.* **72**, 593–607 (2021).
44. Ruamps, L. S. et al. Regulation of soil organic C mineralisation at the pore scale. *FEMS Microbiol. Ecol.* **86**, 26–35 (2013).
45. Barberán, A. et al. Why are some microbes more ubiquitous than others? Predicting the habitat breadth of soil bacteria. *Ecol. Lett.* **17**, 794–802 (2014).
46. Delgado-Baquerizo, M. et al. A global atlas of the dominant bacteria found in soil. *Science* **359**, 320–325 (2018).
47. Juyal, A., Guber, A., Oerther, M., Quigley, M. & Kravchenko, A. Pore architecture and particulate organic matter in soils under monoculture switchgrass and restored prairie in contrasting topography. *Sci. Rep.* **11**, 21998 (2021).
48. Gaillard, V., Chenu, C. & Recous, S. Carbon mineralisation in soil adjacent to plant residues of contrasting biochemical quality. *Soil Biol. Biochem.* **35**, 93–99 (2003).
49. Gaillard, V., Chenu, C., Recous, S. & Richard, G. Carbon, nitrogen and microbial gradients induced by plant residues decomposing in soil. *Eur. J. Soil Sci.* **50**, 567–578 (1999).
50. Kuzyakov, Y. & Blagodatskaya, E. Microbial hotspots and hot moments in soil: concept & review. *Soil Biol. Biochem.* **83**, 184–199 (2015).
51. Ortega-Ramírez, P. et al. Pore distances of particulate organic matter predict N_2O emissions from intact soil at moist conditions. *Geoderma* **429**, 116224 (2023).
52. Kantola, I. B., Masters, M. D. & DeLucia, E. H. Soil particulate organic matter increases under perennial bioenergy crop agriculture. *Soil Biol. Biochem.* **113**, 184–191 (2017).
53. Ulbrich, T. C., Rivas-Ubach, A., Tiemann, L. K., Friesen, M. L. & Evans, S. E. Plant root exudates and rhizosphere bacterial communities shift with neighbor context. *Soil Biol. Biochem.* **172**, 108753 (2022).
54. El Moujahid, L. et al. Effect of plant diversity on the diversity of soil organic compounds. *PLoS ONE* **12**, e0170494 (2017).
55. Harper, C. J., Hayward, D., Kidd, M., Wiid, I. & van Helden, P. Glutamate dehydrogenase and glutamine synthetase are regulated in response to nitrogen availability in *Mycobacterium smegmatis*. *BMC Microbiol.* **10**, 138 (2010).
56. Fazzolari, É., Nicolardot, B. & Germon, J. C. Simultaneous effects of increasing levels of glucose and oxygen partial pressures on denitrification and dissimilatory nitrate reduction to ammonium in repacked soil cores. *Eur. J. Soil Biol.* **34**, 47–52 (1998).
57. Strohm Tobin, O., Griffin, B., Zumft Walter, G. & Schink, B. Growth yields in bacterial denitrification and nitrate ammonification. *Appl. Environ. Micro.* **73**, 1420–1424 (2007).
58. Brewer Tess, E. et al. Ecological and genomic attributes of novel bacterial taxa that thrive in subsurface soil horizons. *Mbio* **10**, e01318–e01319 (2019).
59. Cordero, P. R. F. et al. Carbon monoxide dehydrogenases enhance bacterial survival by oxidising atmospheric CO . Preprint at <https://www.biorxiv.org/content/10.1101/628081v2> (2019).
60. Lucas, M. & Vetterlein, D. X-ray imaging of root–soil interactions. In *X-ray Imaging of the Soil Porous Architecture* (eds Jon Mooney, S. et al.) 129–157 (Springer International Publishing, 2022).
61. Zhou, H. et al. The interaction between wheat roots and soil pores in structured field soil. *J. Exp. Bot.* **72**, 747–756 (2021).
62. Quigley, M. Y. & Kravchenko, A. N. Inputs of root-derived carbon into soil and its losses are associated with pore-size distributions. *Geoderma* **410**, 115667 (2022).
63. Quigley, M. Y., Negassa, W. C., Guber, A. K., Rivers, M. L. & Kravchenko, A. N. Influence of pore characteristics on the fate and distribution of newly added carbon. *Front. Environ. Sci.* **6**, 51 (2018).
64. Thammahong, A., Puttikamonkul, S., Perfect John, R., Brennan Richard, G. & Cramer Robert, A. Central role of the trehalose biosynthesis pathway in the pathogenesis of human fungal infections: opportunities and challenges for therapeutic development. *Microbiol. Mol. Biol. R.* **81**, e00053–00016 (2017).
65. Schimel, J. Modeling ecosystem-scale carbon dynamics in soil: the microbial dimension. *Soil Biol. Biochem.* **178**, 108948 (2023).
66. Slessarev, E. W. & Schimel, J. P. Partitioning sources of CO_2 emission after soil wetting using high-resolution observations and minimal models. *Soil Biol. Biochem.* **143**, 107753 (2020).
67. Erktan, A., Or, D. & Scheu, S. The physical structure of soil: determinant and consequence of trophic interactions. *Soil Biol. Biochem.* **148**, 107876 (2020).
68. Louca, S., Parfrey, L. W. & Doebeli, M. Decoupling function and taxonomy in the global ocean microbiome. *Science* **353**, 1272–1277 (2016).
69. Wright, D. A., Killham, K., Glover, L. A. & Prosser, J. I. Role of pore-size location in determining bacterial-activity during predation by protozoa in soil. *Appl. Environ. Micro.* **61**, 3537–3543 (1995).
70. Fierer, N., Bradford, M. A. & Jackson, R. B. Toward an ecological classification of soil bacteria. *Ecology* **88**, 1354–1364 (2007).
71. Adamczyk, M., Perez-Mon, C., Gunz, S. & Frey, B. Strong shifts in microbial community structure are associated with increased litter input rather than temperature in High Arctic soils. *Soil Biol. Biochem.* **151**, 108054 (2020).
72. Bittman, S., Forge, T. A. & Kowalenko, C. G. Responses of the bacterial and fungal biomass in a grassland soil to multi-year applications of dairy manure slurry and fertilizer. *Soil Biol. Biochem.* **37**, 613–623 (2005).
73. Kleber, M. & Lindsley, A. The science and semantics of “Soil organic matter stabilization”. In *Multi-Scale Biogeochemical Processes in Soil Ecosystems*. (eds Yang, Y., Keiluweit, M., Senesi, N. & Xing, B.) 13–49 (John Wiley & Sons, Inc., 2022).
74. Pot, V. et al. Understanding the joint impacts of soil architecture and microbial dynamics on soil functions: insights derived from microscale models. *Eur. J. Soil Sci.* **73**, e13256 (2022).
75. Sulman, B. N., Phillips, R. P., Oishi, A. C., Shevliakova, E. & Pacala, S. W. Microbe-driven turnover offsets mineral-mediated storage of soil carbon under elevated CO_2 . *Nat. Clim. Change* **4**, 1099–1102 (2014).
76. Robertson, G. P. & Hamilton, S. K. Long-term ecological research in agricultural landscapes at the Kellogg Biological Station LTER site: conceptual and experimental framework. In *The Ecology of Agricultural Landscapes: Long-term Research on the Path to Sustainability* (eds Hamilton, S. K. et al.) 1–32 (Oxford University Press, 2015).
77. Lee, J. H., Lucas, M., Guber, A. K., Li, X. & Kravchenko, A. N. Interactions among soil texture, pore structure, and labile carbon influence soil carbon gains. *Geoderma* **439**, 116675 (2023).
78. Nelson, D., Sommers, L. & Sparks, D. Method of soil analysis: chemical methods. *Methods Soil Anal. Part 3 Chem. Methods* **3**, 961–1010 (1996).
79. Thomas, G. W. Soil pH and soil acidity. In *Methods of Soil Analysis*. (eds Sparks, D. L. et al.) 475–490 (Soil Science Society of America, Inc., American Society of Agronomy, Inc., 1996).
80. Kuo, S. Phosphorus. In *Methods of Soil Analysis*. (eds Sparks, D. L. et al.) 869–919 (Soil Science Society of America, Inc., American Society of Agronomy, Inc., 1996).
81. Sumner M. E. & Miller, W. P. Cation exchange capacity and exchange coefficients. In *Methods of Soil Analysis*. (eds Sparks, D. L. et al.) 1201–1229 (Soil Science Society of America, Inc., American Society of Agronomy, Inc., 1996).

82. Lauenroth, W. K. Methods of estimating belowground net primary production. In: *Methods in Ecosystem Science* (eds Sala, O. E. et al.) 58–71 (Springer New York, 2000).
83. Grossman, R. B. & Reinsch, T. G. 2.1 Bulk density and linear extensibility. In *Methods of Soil Analysis*. (eds Dane, J. H. & Topp, G. C.) 201–228 (Soil Science Society of America, Inc., SSSA Book Series, 2002).
84. Chen, Y. & Murrell, J. C. When metagenomics meets stable-isotope probing: progress and perspectives. *Trends Microbiol.* **18**, 157–163 (2010).
85. Robertson, G. P., Coleman, D. C., Bledsoe, C. S. & Sollins, P. *Standard Soil Methods for Long-term Ecological Research* (Oxford University Press, 1999).
86. Vance, E. D., Brookes, P. C. & Jenkinson, D. S. An extraction method for measuring soil microbial biomass C. *Soil Biol. Biochem.* **19**, 703–707 (1987).
87. Rasband, W. S. *ImageJ* (U.S. National Institutes of Health, 1997–2015).
88. Schindelin, J. et al. Fiji: an open-source platform for biological-image analysis. *Nat. Methods* **9**, 676–682 (2012).
89. Kravchenko, A. N., Negassa, W., Guber, A. K. & Schmidt, S. New approach to measure soil particulate organic matter in intact samples using X-ray computed microtomography. *Soil Sci. Soc. Am. J.* **78**, 1177–1185 (2014).
90. Lucas, M., Nguyen, L. T. T., Guber, A. & Kravchenko, A. Cover crop influence on pore size distribution and biopore dynamics: enumerating root and soil faunal effects. *Front. Plant Sci.-Funct. Plant Ecol.* **13**, 928569 (2022).
91. Berg, S. et al. ilastik: interactive machine learning for (bio)image analysis. *Nat. Methods* **16**, 1226–1232 (2019).
92. Doube, M. et al. BoneJ Free and extensible bone image analysis in ImageJ. *Bone* **47**, 1076–1079 (2010).
93. Schluter, S., Weller, U. & Vogel, H. J. Soil-structure development including seasonal dynamics in a long-term fertilization experiment. *J. Plant Nutr. Soil Sci.* **174**, 395–403 (2011).
94. Vogel, H. J., Weller, U. & Schluter, S. Quantification of soil structure based on Minkowski functions. *Comput. Geosci.* **36**, 1236–1245 (2010).
95. Munch, B. & Holzer, L. Contradicting geometrical concepts in pore size analysis attained with electron microscopy and mercury intrusion. *J. Am. Ceram. Soc.* **91**, 4059–4067 (2008).
96. Kravchenko, A., Falconer, R. E., Grinev, D. & Otten, W. Fungal colonization in soils with different management histories: modeling growth in three-dimensional pore volumes. *Ecol. Appl.* **21**, 1202–1210 (2011).
97. Kandeler, E., Luxhøi, J., Tscherko, D. & Magid, J. Xylanase, invertase and protease at the soil–litter interface of a loamy sand. *Soil Biol. Biochem.* **31**, 1171–1179 (1999).
98. Védère, C., Vieublé Gonod, L., Pouteau, V., Girardin, C. & Chenu, C. Spatial and temporal evolution of detritusphere hotspots at different soil moistures. *Soil Biol. Biochem.* **150**, 107975 (2020).
99. Kuz'yakov, Y. & Razavi, B. S. Rhizosphere size and shape: temporal dynamics and spatial stationarity. *Soil Biol. Biochem.* **135**, 343–360 (2019).
100. Moran, J. J. et al. Spectroscopy-based isotopic ($\delta^{13}\text{C}$) analysis for high spatial resolution of carbon exchange in the rhizosphere. *Rhizosphere* **23**, 100564 (2022).
101. Kim, K., Kaestner, A., Lucas, M. & Kravchenko, A. N. Microscale spatiotemporal patterns of water, soil organic carbon, and enzymes in plant litter detritusphere. *Geoderma* **438**, 116625 (2023).
102. Kravchenko, A. N. et al. Hotspots of soil N_2O emission enhanced through water absorption by plant residue. *Nat. Geosci.* **10**, 496 (2017).
103. Kutlu, T., Guber, A. K., Rivers, M. L. & Kravchenko, A. N. Moisture absorption by plant residue in soil. *Geoderma* **316**, 47–55 (2018).
104. Guber, A., Kutlu, T., Rivers, M. & Kravchenko, A. Mass-balance approach to quantify water distribution in soils based on X-ray computed tomography images. *Eur. J. Soil Sci.* **72**, 578–592 (2021).
105. Crawford, J. W. et al. Microbial diversity affects self-organization of the soil-microbe system with consequences for function. *J. R. Soc. Interface* **9**, 1302–1310 (2012).
106. Hoshino, Y. T. & Naoyuki, M. An improved DNA extraction method using skim milk from soils that strongly adsorb DNA. *Microbes Environ.* **19**, 13–19 (2004).
107. Kozich, J. J., Westcott, S. L., Baxter, N. T., Highlander, S. K. & Schloss, P. D. Development of a dual-index sequencing strategy and curation pipeline for analyzing amplicon sequence data on the miseq illumina sequencing platform. *Appl. Environ. Micro.* **79**, 5112–5120 (2013).
108. Schloss, P. D. et al. Introducing Mothur: open-source, platform-independent, community-supported software for describing and comparing microbial communities. *Appl. Environ. Micro.* **75**, 7537–7541 (2009).
109. Pruesse, E. et al. SILVA: a comprehensive online resource for quality checked and aligned ribosomal RNA sequence data compatible with ARB. *Nucleic Acids Res.* **35**, 7188–7196 (2007).
110. R_Core_Team. R: A Language and Environment for Statistical Computing. (Vennia, Auatria, 2020).
111. RStudio_Team. RStudio: Integrated Development for R. RStudio, PBC. (Boston, MA, 2020).
112. McMurdie, P. J. & Holmes, S. phyloseq: an R package for reproducible interactive analysis and graphics of microbiome census data. *PLoS ONE* **8**, e61217 (2013).
113. Liu, C., Cui, Y., Li, X. & Yao, M. microeco: an R package for data mining in microbial community ecology. *FEMS Microbiol. Ecol.* **97**, fiaa255 (2021).
114. Lahti, L. & Shetty, S. Microbiome R package. (2012–2019).
115. Andersen, K. S., Kirkegaard, R. H., Karst, S. M. & Albertsen, M. ampvis2: an R package to analyse and visualise 16 S rRNA amplicon data. Preprint at <https://www.biorxiv.org/content/10.1101/299537v1.full.pdf> (2018).
116. Oksanen, J. et al. Vegan: Community Ecology Package. (2020).
117. Martinez, A. P. pairwiseAdonis: pairwise multilevel comparison using Adonis. R package version 0.4. (2017).
118. Gu, Z., Gu, L., Eils, R., Schlesner, M. & Brors, B. Circlize implements and enhances circular visualization in R. *Bioinformatics* **30**, 2811–2812 (2014).
119. Wickham, H., François, R., Henry, L., Müller, K. & Vaughan, D. *dplyr: A Grammar of Data Manipulation*. R package version 1.1.4, <https://github.com/tidyverse/dplyr>, <https://dplyr.tidyverse.org> (2023).
120. Wickham, H., Vaughan, D. & Girlich, M. *tidyr: Tidy Messy Data*. R package version 1.3.1, <https://github.com/tidyverse/tidyr>, <https://tidyr.tidyverse.org> (2024).
121. Kassambara, A. *ggpubr: 'ggplot2' Based Publication Ready Plots*. R package version 0.6.0, <https://rpkgs.datanovia.com/ggpubr/> (2023).
122. Kassambara, A. *rstatix: Pipe-Friendly Framework for Basic Statistical Tests*. R package version 0.7.2, <https://rpkgs.datanovia.com/rstatix/> (2023).
123. Wickham, H. *ggplot2: Create Elegant Data Visualisations Using the Grammar of Graphics* (Springer-Verlag New York, 2016).
124. Douglas, G. M. et al. PICRUSt2 for prediction of metagenome functions. *Nat. Biotechnol.* **38**, 685–688 (2020).
125. Kanehisa, M., Sato, Y., Kawashima, M., Furumichi, M. & Tanabe, M. KEGG as a reference resource for gene and protein annotation. *Nucleic Acids Res.* **44**, D457–D462 (2016).

126. Barbera, P. et al. EPA-ng: massively parallel evolutionary placement of genetic sequences. *Syst. Biol.* **68**, 365–369 (2019).
 127. Czech, L., Barbera, P. & Stamatakis, A. Genesis and Gappa: processing, analyzing and visualizing phylogenetic (placement) data. *Bioinformatics* **36**, 3263–3265 (2020).
 128. Louca, S. & Doebeli, M. Efficient comparative phylogenetics on large trees. *Bioinformatics* **34**, 1053–1055 (2017).
 129. Ye, Y. & Doak, T. G. A parsimony approach to biological pathway reconstruction/inference for genomes and metagenomes. *PLoS Comput. Biol.* **5**, e1000465 (2009).
 130. Milliken, G. A. & Johnson, D. E. *Analysis of Messy Data Volume I: Designed Experiments*, 2nd edn. (CRC Press, 2009).
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Competing interests

The authors declare no competing interests.

Additional information

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Author contributions

Z.L. conducted SIP experiment and analyses of biological data; A.N.K., A.K.G., and A.C. developed research concepts with inputs from G.P.R.,